SOUTH AMERICA: A RESERVOIR OF CONTINENTAL CARBON - FIRST ESTIMATE OF CHANGES SINCE 18,000 YR BP

ABSTRACT

By using geographic and palaeogeographic sketches established for the present situation (before recent deforestation) and for the glacial maximum (about 15,000-18,000 BP) we can estimate the possible total biomass (phytomass) of the South American continent. According to the biomass density used in this first estimate for ten major ecosystems, the results show a possible increase from 140 Gt of carbon (glacial maximum) to 214 Gt C (preindustrial) for the phytomass, and 120 to 180 Gt C for the soils.

These preliminary results are possibly only a 60 or 70 percent approximate estimate and could be modified with computation using other palaeogeographic models or another biomass density. It is

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therefore to underline the urgent need of more field biomass measurements, ecosystems mappings, and palaeostudies to evaluate the part of South America as a future possible sink for the atmospheric carbon dioxide.

The Amazonian forest makes of South America an important continental reservoir of carbon for the planet Earth.

This continent represents consequently a key zone for the research and knowledge of changes in the biogeochemical cycle of carbon. In order to evaluate more precisely the role it plays we estimated the approximate quantities of carbon in the total phytomass and the carbon in soils for each of the ecosystems represented in Figure 1, both for Present and Last Glacial Maximum landscapes.

METHOD

For first preliminary estimate a sketch map of the eight principal biomes or ecosystems of South America was organized (FAURE et al., 1985) based upon the palaeobioclimatic maps by HUECK (1972), AB'SABER (1977), BROWN & AB'SABER (1979), and palaeopedological and geochemical maps by PEDRO & VOLKOFF (1984). Two representation are given: one for the Present state (before the great deforestation), and another for a probable state during the end of the Last Glacial Maximum (18,000-15,000 yr BP), according to palaeoclimatic data of the authors. Areas of principal ecosystems were measured and the most probable biomass value was calculated for each ecosystem using the mean density values compiled for the Earth by AJJAY et al. (1979). These values are closer to more recent field measurements (MAURY-LECHON, 1982) and slightly higher than those from volumetric evaluation (BROWN & LUGO, 1984).

RESULTS

Results in Table I allow to propose a first estimate of total biomass and organic matter of soils for South America, calculated in carbon, for two opposite climatic states. The lack of precision about the Present areas occupied by the ecosystems and their mean biomass causes an standard deviation of 30-40%. Considering the palaeogeographic documents and densities of carbon/m² we used for this preliminary estimate, the model for the Present situation gives a value of 214 Gt (Gt=Gigaton=10^{15} g) of vegetal biomass, and 140 Gt for the Glacial Maximum; reduction of vegetal carbon would amount to 1/3 during a glacial phase. It would be useful to recalculate these data with other palaeogeographic models and other biomass values to obtain a more precise standard deviation.
Figure 1 - Simplified map of the principal ecosystems in South America for a Glacial Maximum and for "Present".
### Table 1 - South America biomass. From left to right: ecosystem, carbon density/m² (biomass in the soil upper meter), area (millions of km²), total biomass, total weight of soil carbon (carbon Gt) for Present and for 18,000BP.

<table>
<thead>
<tr>
<th>Ecosystem</th>
<th>Biomass Ckg/m²</th>
<th>Soils Mkm²</th>
<th>Present situation Biomass Mkm²</th>
<th>Soils CGt</th>
<th>Glacial situation (18,000BP) Biomass Mkm²</th>
<th>Soils CGt</th>
</tr>
</thead>
<tbody>
<tr>
<td>Humid forests</td>
<td>20</td>
<td>15</td>
<td>6.4</td>
<td>128.0</td>
<td>96.0</td>
<td>1.7</td>
</tr>
<tr>
<td>Humid forests (coasts)</td>
<td>20</td>
<td>15</td>
<td>0.4</td>
<td>8.0</td>
<td>6.0</td>
<td>0.2</td>
</tr>
<tr>
<td>Cerrados</td>
<td>15</td>
<td>7</td>
<td>3.7</td>
<td>55.5</td>
<td>25.9</td>
<td>5.3</td>
</tr>
<tr>
<td>Araucaria forests</td>
<td>10</td>
<td>15</td>
<td>1</td>
<td>10.0</td>
<td>15.0</td>
<td>0.9</td>
</tr>
<tr>
<td>Caatinga</td>
<td>4</td>
<td>7</td>
<td>0.8</td>
<td>2.4</td>
<td>5.6</td>
<td>4.0</td>
</tr>
<tr>
<td>Herbaceous steppes</td>
<td>2</td>
<td>12</td>
<td>1</td>
<td>2.0</td>
<td>12.0</td>
<td>0.0</td>
</tr>
<tr>
<td>Subdesertic steppes</td>
<td>0.3</td>
<td>6</td>
<td>0.9</td>
<td>0.3</td>
<td>5.4</td>
<td>1.3</td>
</tr>
<tr>
<td>Desertic steppes</td>
<td>0.01</td>
<td>1</td>
<td>0.8</td>
<td>0.0</td>
<td>0.8</td>
<td>1.1</td>
</tr>
<tr>
<td>Flooded, etc.</td>
<td>8</td>
<td>10</td>
<td>0.9</td>
<td>7.2</td>
<td>9.0</td>
<td>0.2</td>
</tr>
<tr>
<td>Mountains, coastal deposits</td>
<td>0.1</td>
<td>0.5</td>
<td>2</td>
<td>0.2</td>
<td>1.0</td>
<td>3.8</td>
</tr>
<tr>
<td>Total</td>
<td></td>
<td></td>
<td>17.9</td>
<td>213.6</td>
<td>176.7</td>
<td>18.5</td>
</tr>
</tbody>
</table>

About soils as carbon reservoirs a first continental estimate may be proposed. Considering timescales between 100-1,000 years it could be said that soil carbon content between 0-100 cm beneath the surface balances with the vegetation supported by that soil. From SANCHEZ et al. (1982) data and measures by CERRI & VOLKOFF (1987) the carbon stock for the upper meter of soil has been calculated for each ecosystem. Table I shows that this stock represents today about 180 Gt; during the Last Glacial Maximum its value was probably about 120 Gt. A fluctuation of about 60 Gt could be estimated for this soil reservoir.

### COMPARISON

Notwithstanding their preliminary character, these results can be compared to a previous estimate of the world continental biomass carbon (FAURE et al., 1989). Figures were calculated with a 30% estimated error for the world 1975 continental biomass, based upon AJJAY et al. (1979) values (Table II). The increase of the continental vegetation carbon reservoir from Last Glacial Maximum to Present would be of about 268 Gt (48 of Present); South American contribution would be about 73 Gt.

### Table 2 - Changes in the carbon reservoir of continental vegetation (MGT of carbon).

<table>
<thead>
<tr>
<th></th>
<th>18,000BP</th>
<th>Present (1975 A.D.)</th>
</tr>
</thead>
<tbody>
<tr>
<td>Total forest phytomass</td>
<td>106</td>
<td>428</td>
</tr>
<tr>
<td>Total non forest phytomass</td>
<td>186</td>
<td>132</td>
</tr>
<tr>
<td>Total</td>
<td>292</td>
<td>560</td>
</tr>
</tbody>
</table>

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CONCLUSIONS

These approximations show important variations and should be improved. Therefore it is necessary to augment biomass and soil carbon evaluations; to continue with ecosystems cartography and dated palaeoclimatic studies in order to corroborate or to reconsider the role played by South America as a carbon reservoir, as well as the role it may play in the future to contribute to the maintenance of an adequate atmosphere for Man (Fig. 2). For the present preliminary results suggest that the glacial-to-interglacial carbon dioxide increase of about 200 Gt (BARNOLA et al., 1987) was accompanied by a similar or greater increase in the biosphere, and probably doubled by the soil carbon pool increase. In this global continental carbon increase South American contribution could be close to about 129 Gt (in carbon).

Figure 2 - There is an urgent need for maps giving the distributions of major ecotypes at precise dates in the Quaternary (from HENDERSON-SELLERS et al., 1988).
REFERENCES


